

紀伊水道における海洋短波レーダーを 利用した津波早期警報

Tsunami early warning using high-frequency ocean radar system
in the Kii Channel, Japan

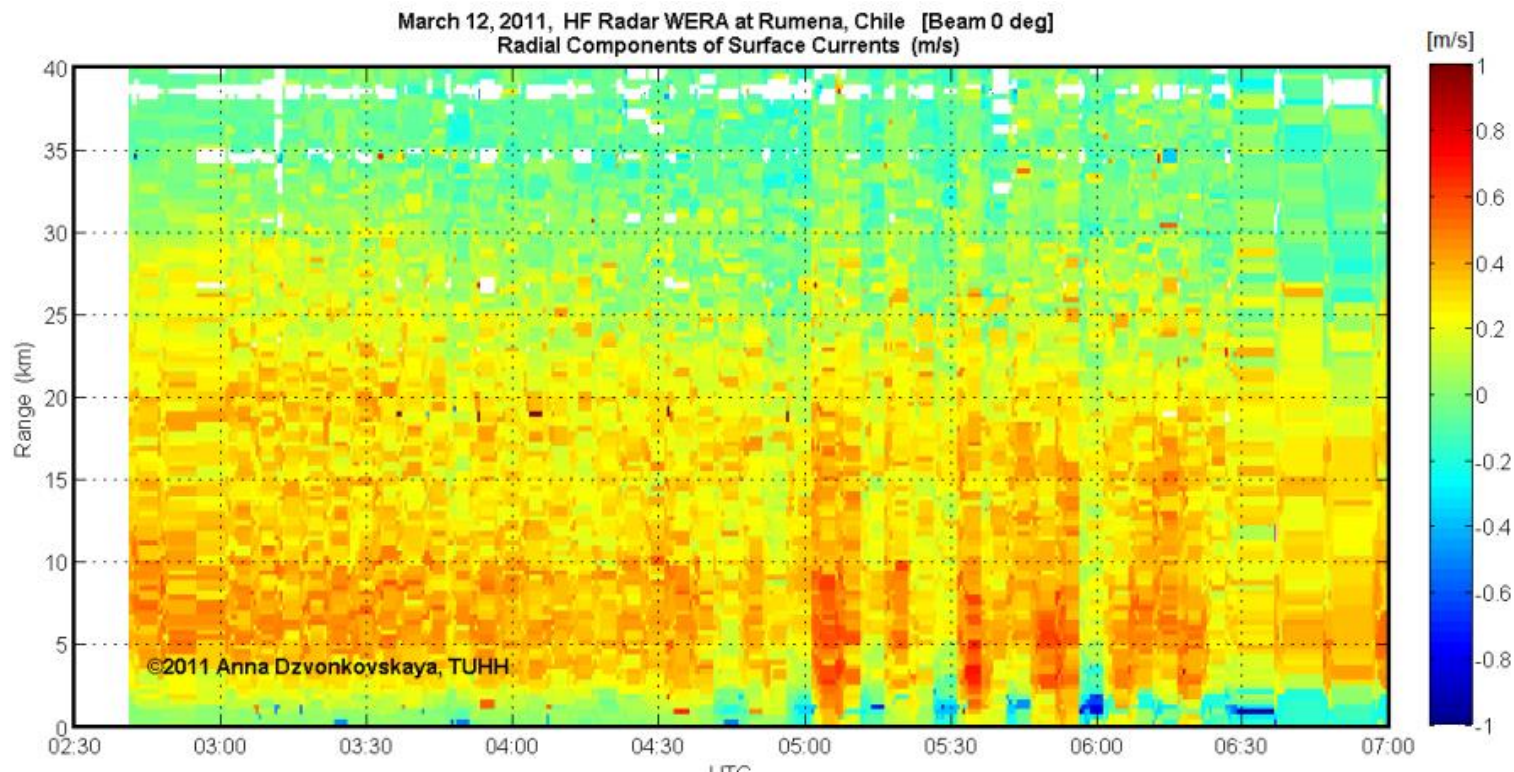
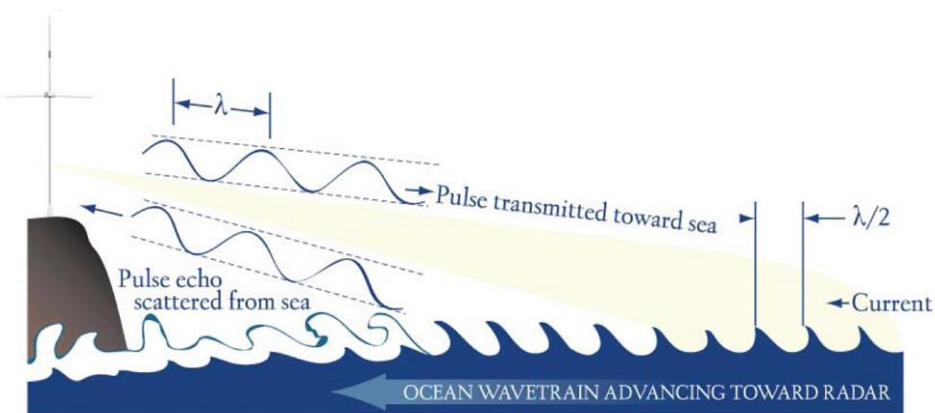
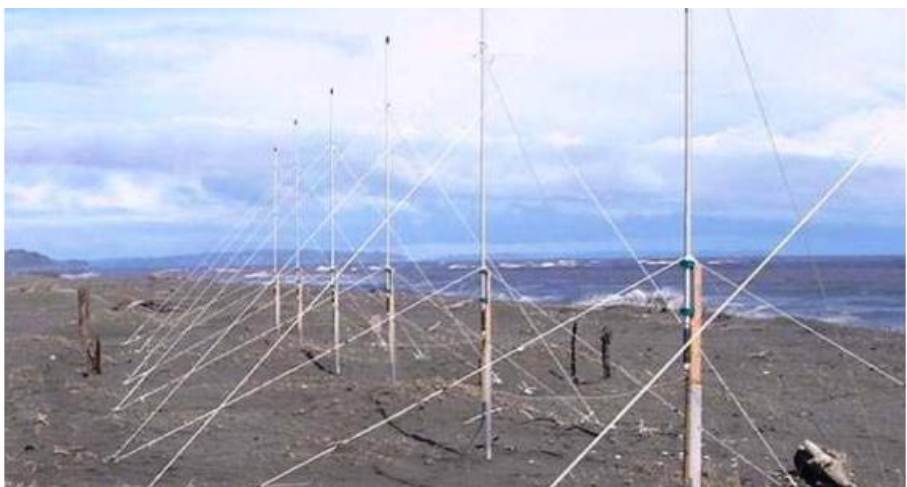
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国立研究開発法人海洋研究開発機構

1. はじめに

High-frequency (HF) ocean radar

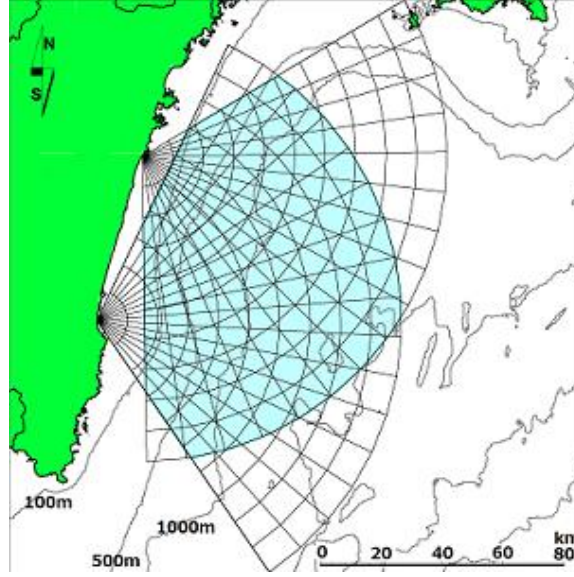
Example: HF radar in Chile observed the tsunami of the 2011 Tohoku earthquake.



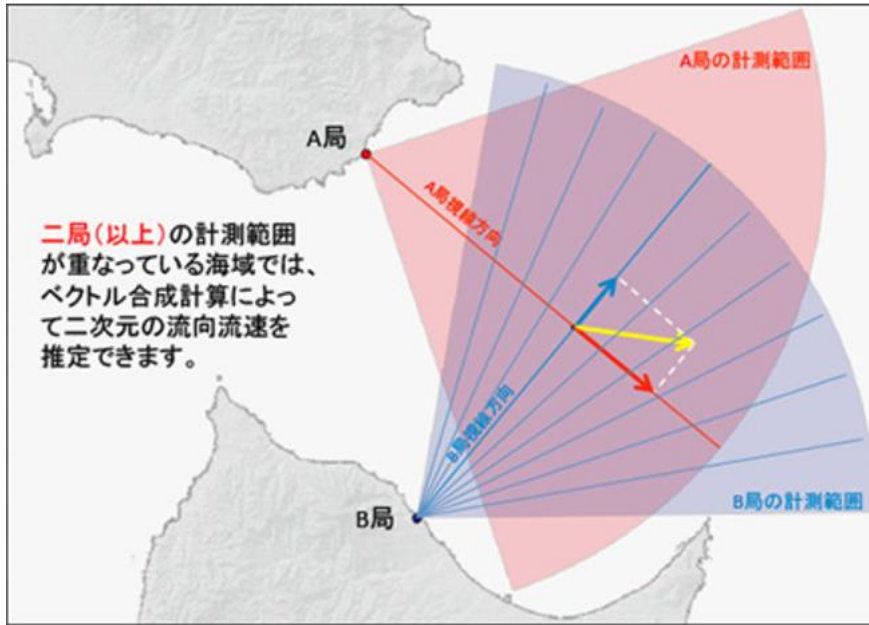
(Dzvonkovskaya et al., [2011](#); Roarty et al., [2019](#))

HF radar in Japan

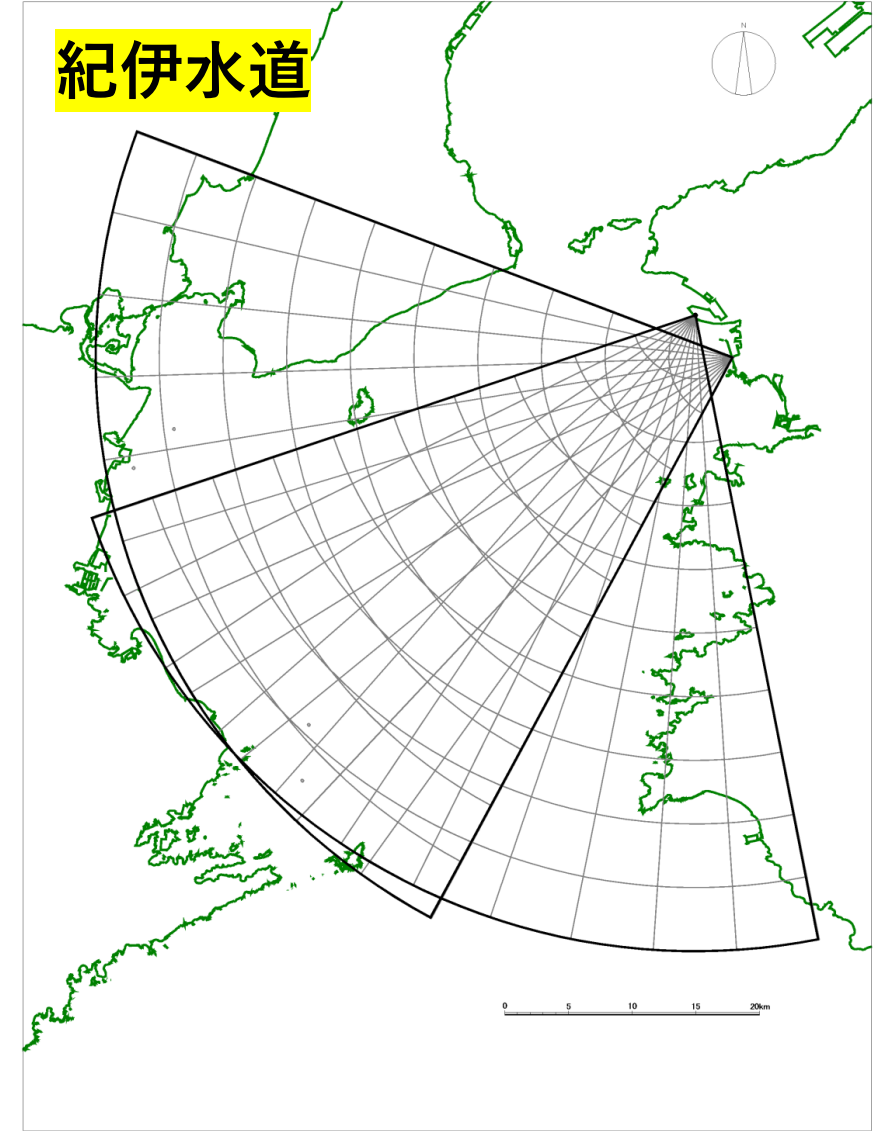
日向灘



津軽海峡



紀伊水道



(Figure Source: JAMSTEC, MLIT, Miyazaki Prefecture)

2. データ解析

2011 Tohoku Tsunami

- Generated by the M9 Tohoku Earthquake on March 11, 2011
- Arrived in Kii Channel approximately 3 hours after the earthquake



(Figure Source: The New York Times)



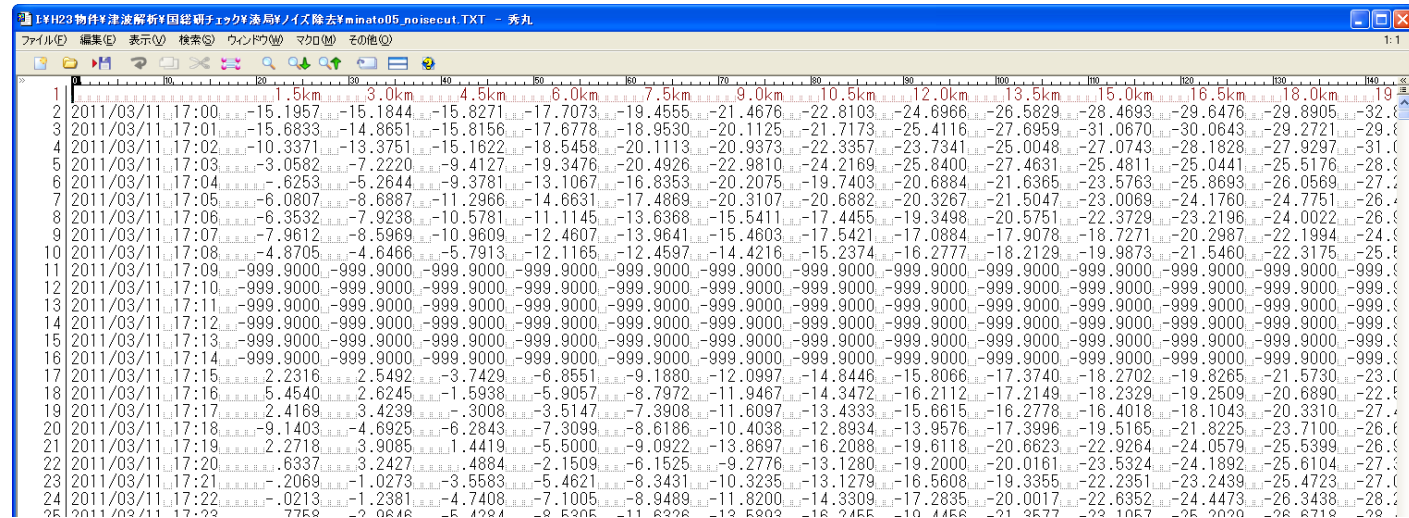
Minato 湊

Saikazaki 雑賀崎

観測データ@湊 (処理前)

The tsunamis of the 2011 Tohoku earthquake was observed by high-frequency radar system in Kii Channel.

Sampling rate: 1 min



134°

135°

136°

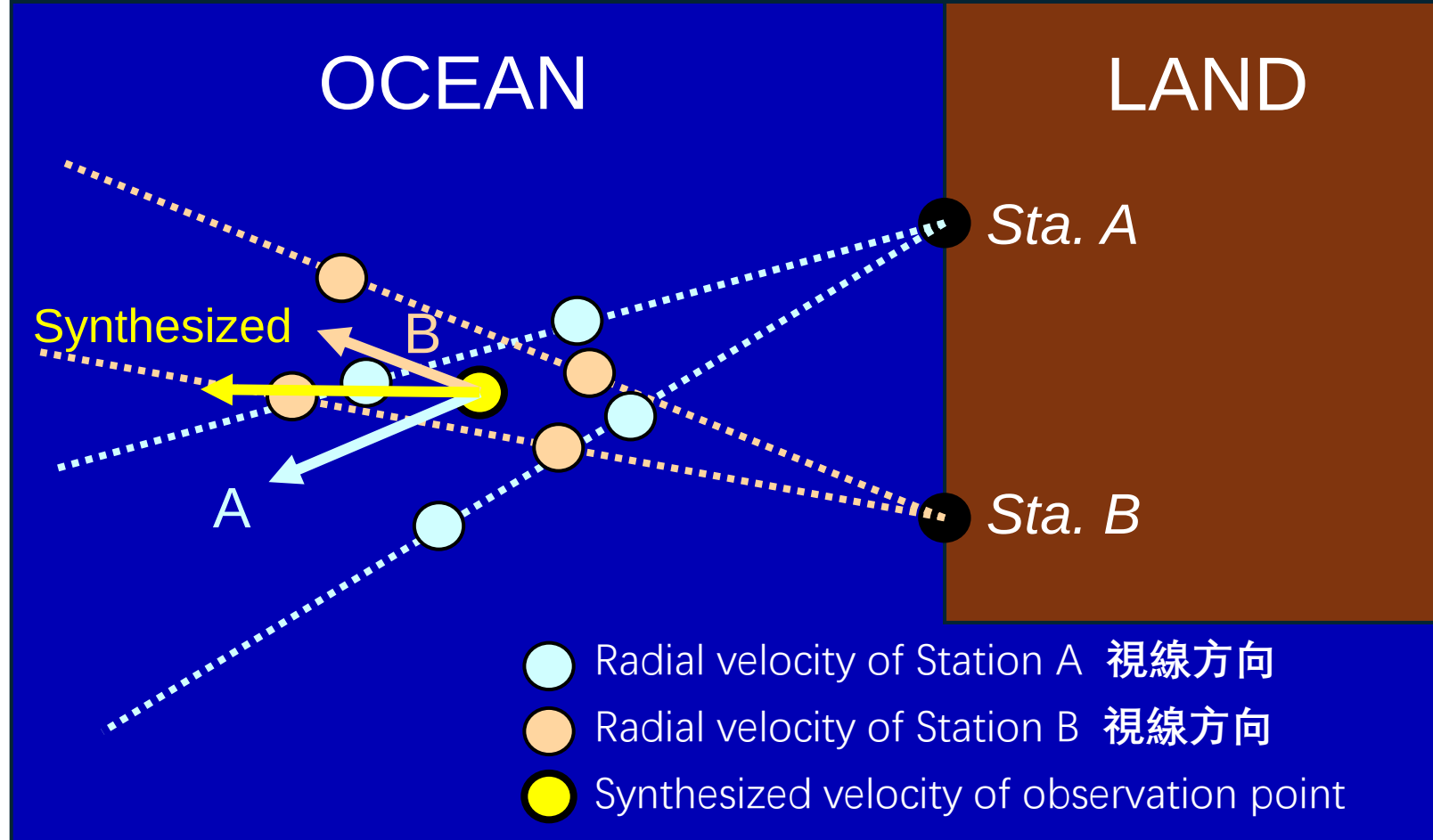
- HF Radar Land Station
- HF Radar Observational Point
- ▲ Tide Gauge

Kobe*OSAKA BAY***Minato sta.****Saikazaki sta.***SEA OF JAPAN***Study region***PACIFIC OCEAN**KII
CHANNEL*

10 km

34°

Vector Synthesis of Surface Vector



- At least **two** land stations are required for vector synthesis
- It is necessary to install the beams of the two stations to be close to **90 degrees**

Beam direction (radial velocity) -> NS/EW direction

Vector Synthesis of Surface Vector

$$\begin{bmatrix} \sin\theta_1 & \cos\theta_1 \\ \sin\theta_2 & \cos\theta_2 \\ \vdots & \vdots \\ \sin\theta_N & \cos\theta_N \end{bmatrix} \begin{pmatrix} u \\ v \end{pmatrix} = \begin{bmatrix} U_{r1} \\ U_{r2} \\ \vdots \\ U_{rN} \end{bmatrix}$$

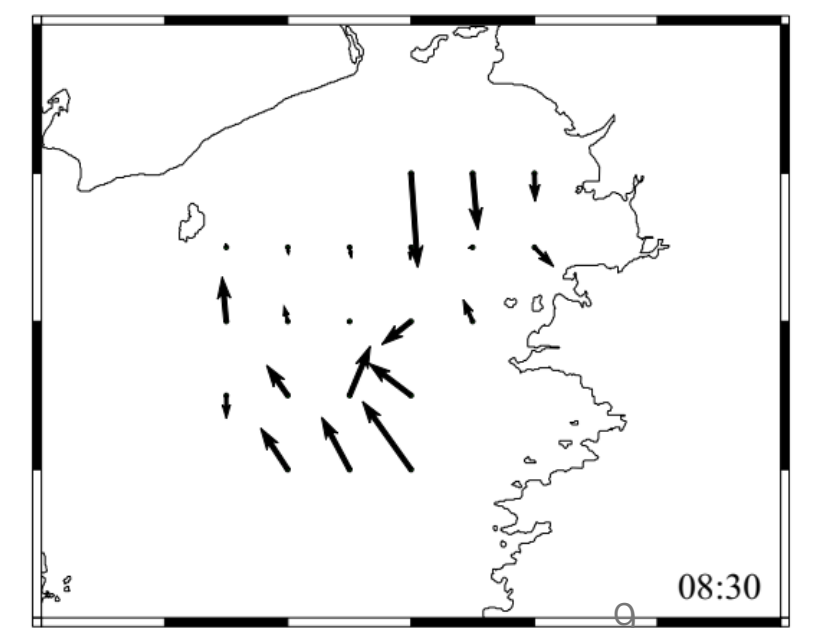
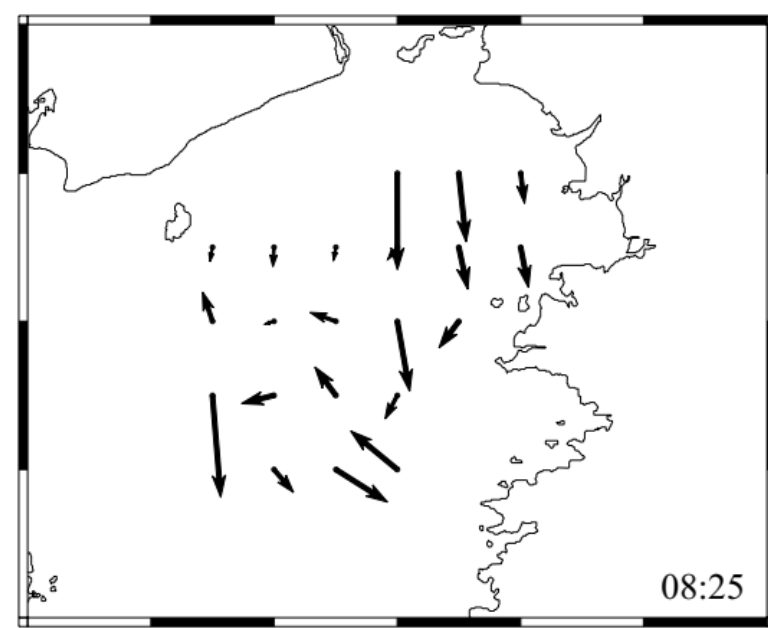
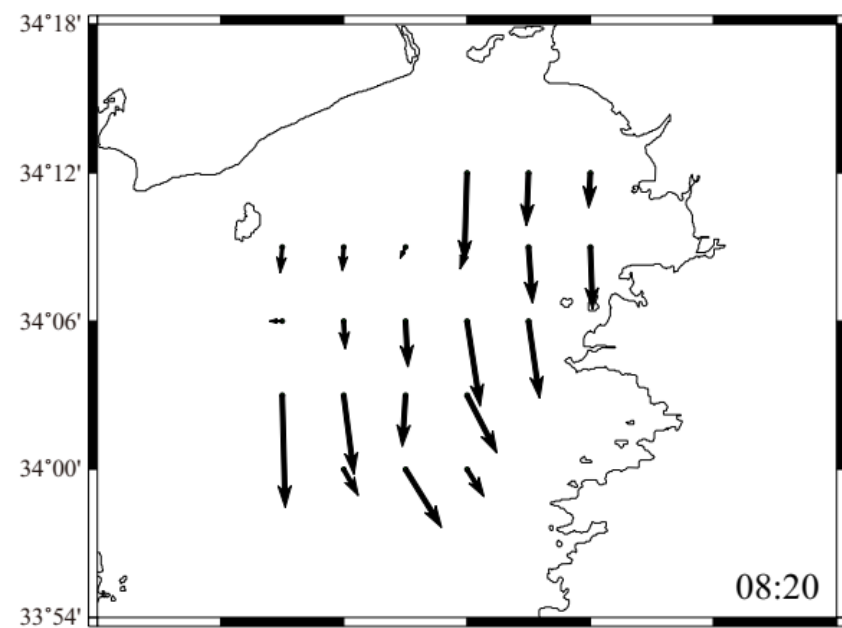
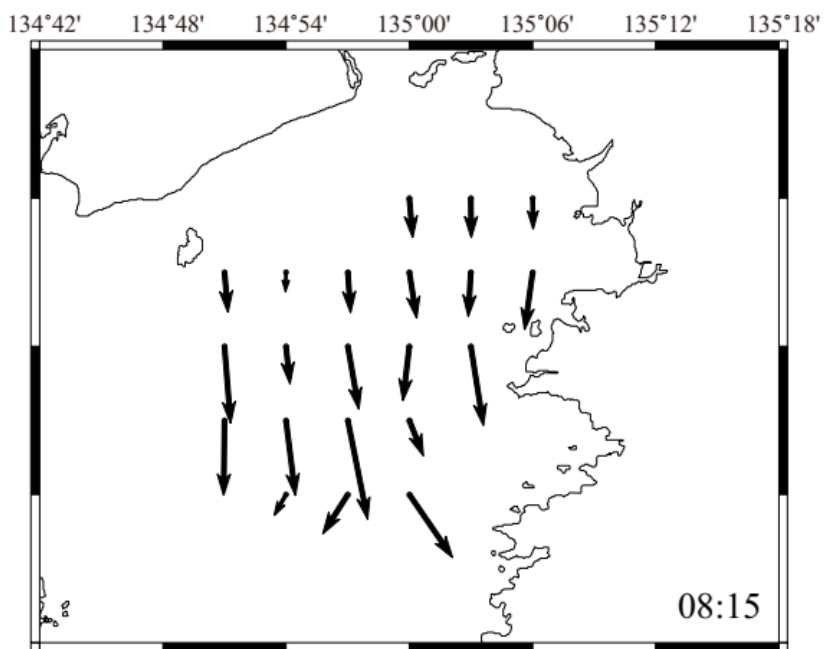
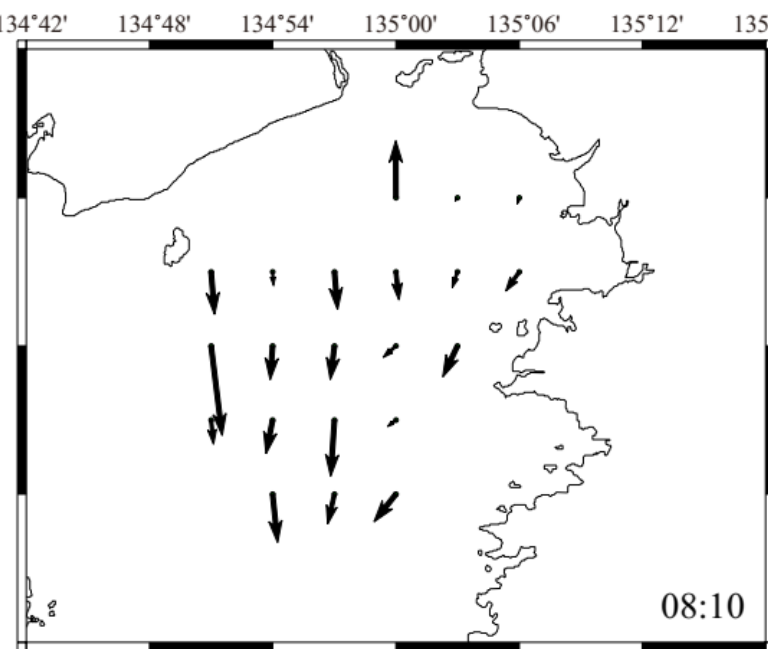
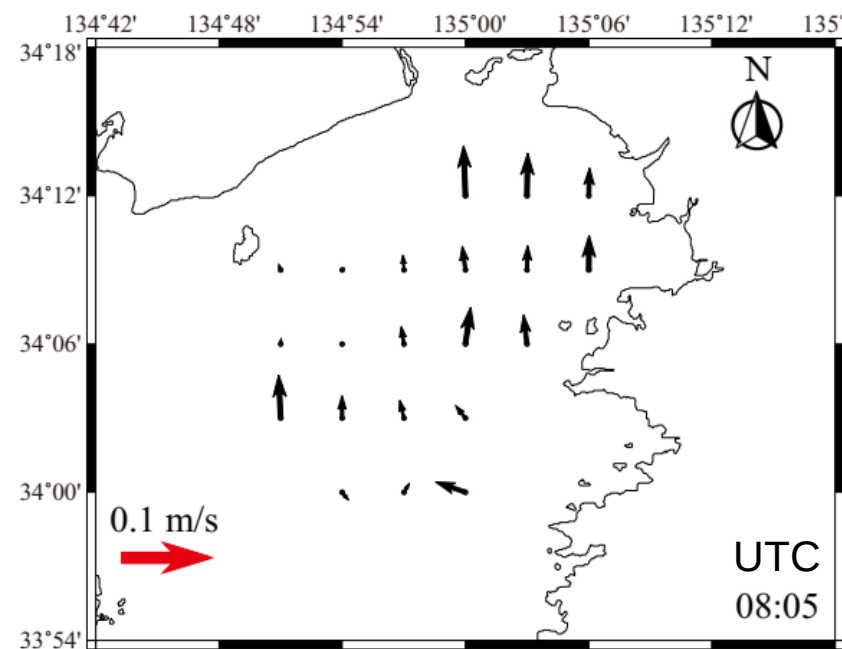
i : Number of HF radar land station
 θ_i : Azimuth of the grid
 U_{ri} : Velocity in the beam direction
 u, v : Velocity in the EW/NS direction

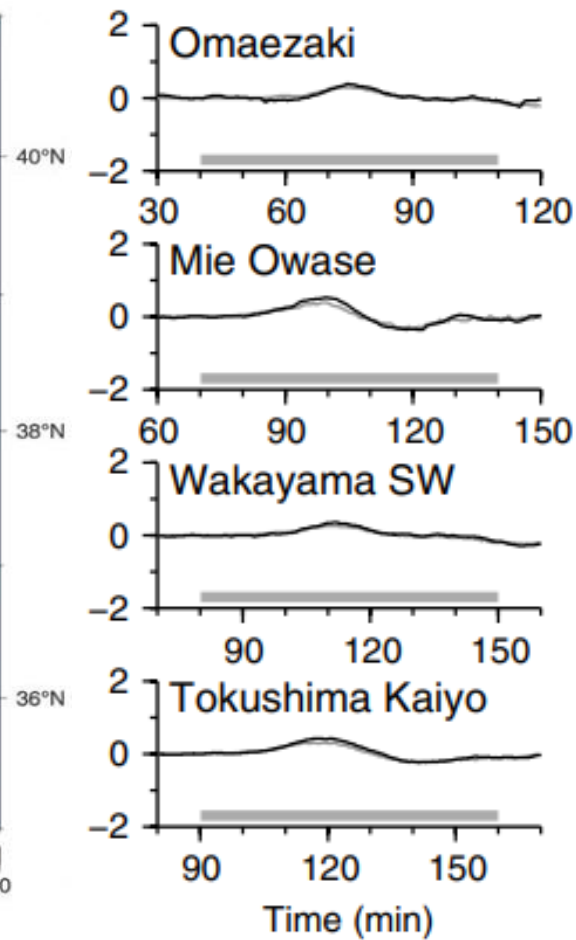
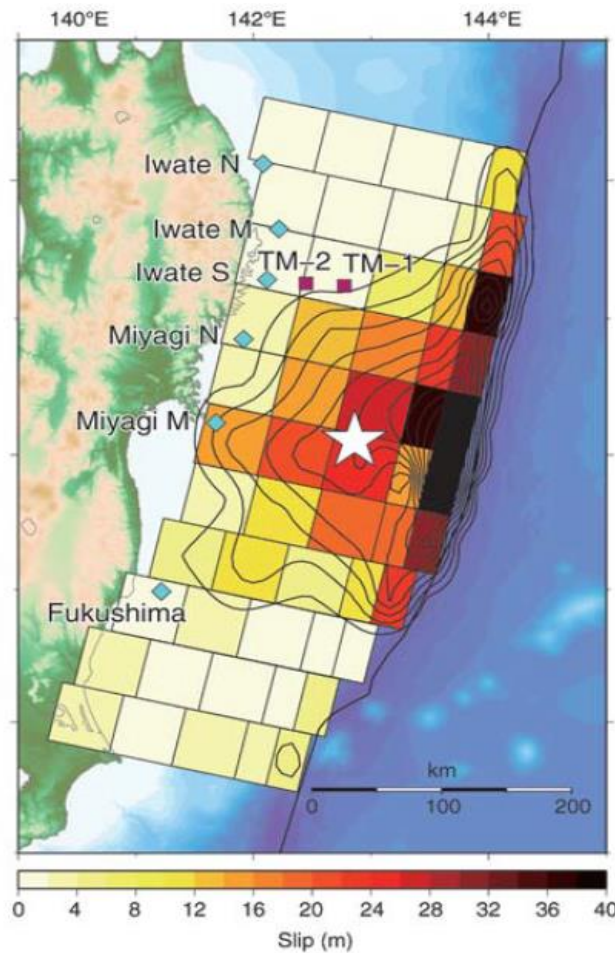
When $N > 2$:

$$\sum_{i=1}^N (U_{ri} - \hat{U}_{r1})^2 / \sigma_i^2 \rightarrow \min$$

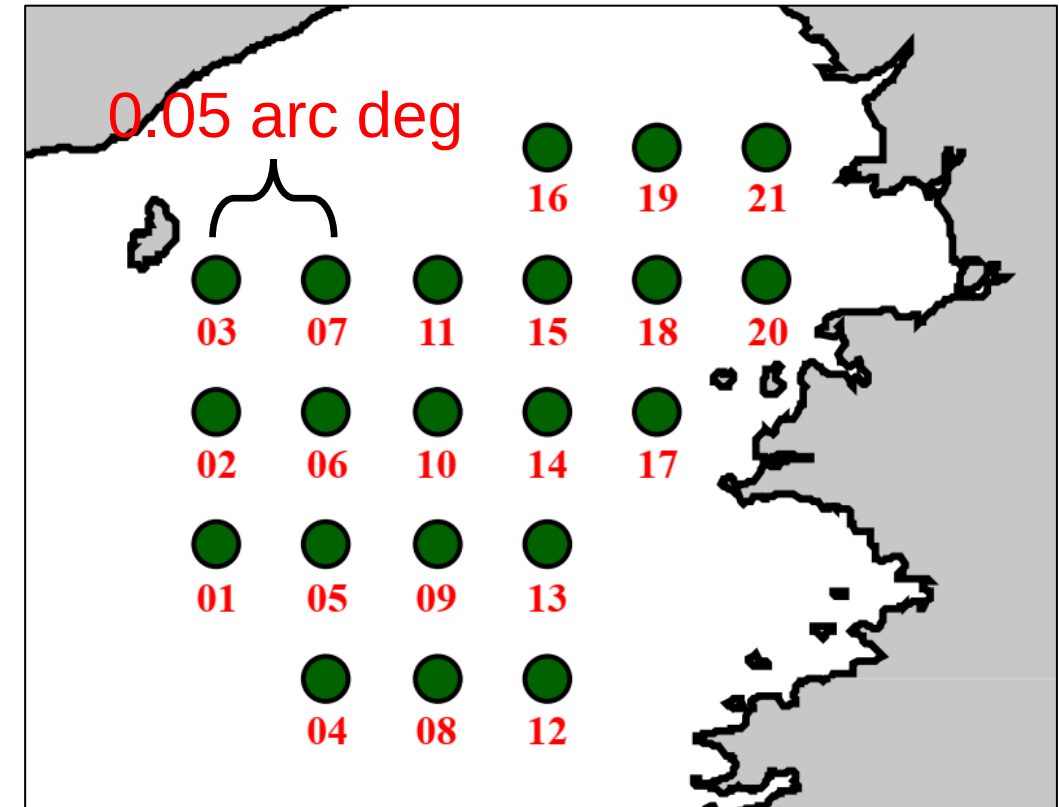
When $N = 2$ (in this study):

$$\begin{pmatrix} u \\ v \end{pmatrix} = \begin{bmatrix} \sin\theta_1 & \cos\theta_1 \\ \sin\theta_2 & \cos\theta_2 \end{bmatrix}^{-1} \begin{bmatrix} U_{r1} \\ U_{r2} \end{bmatrix}$$





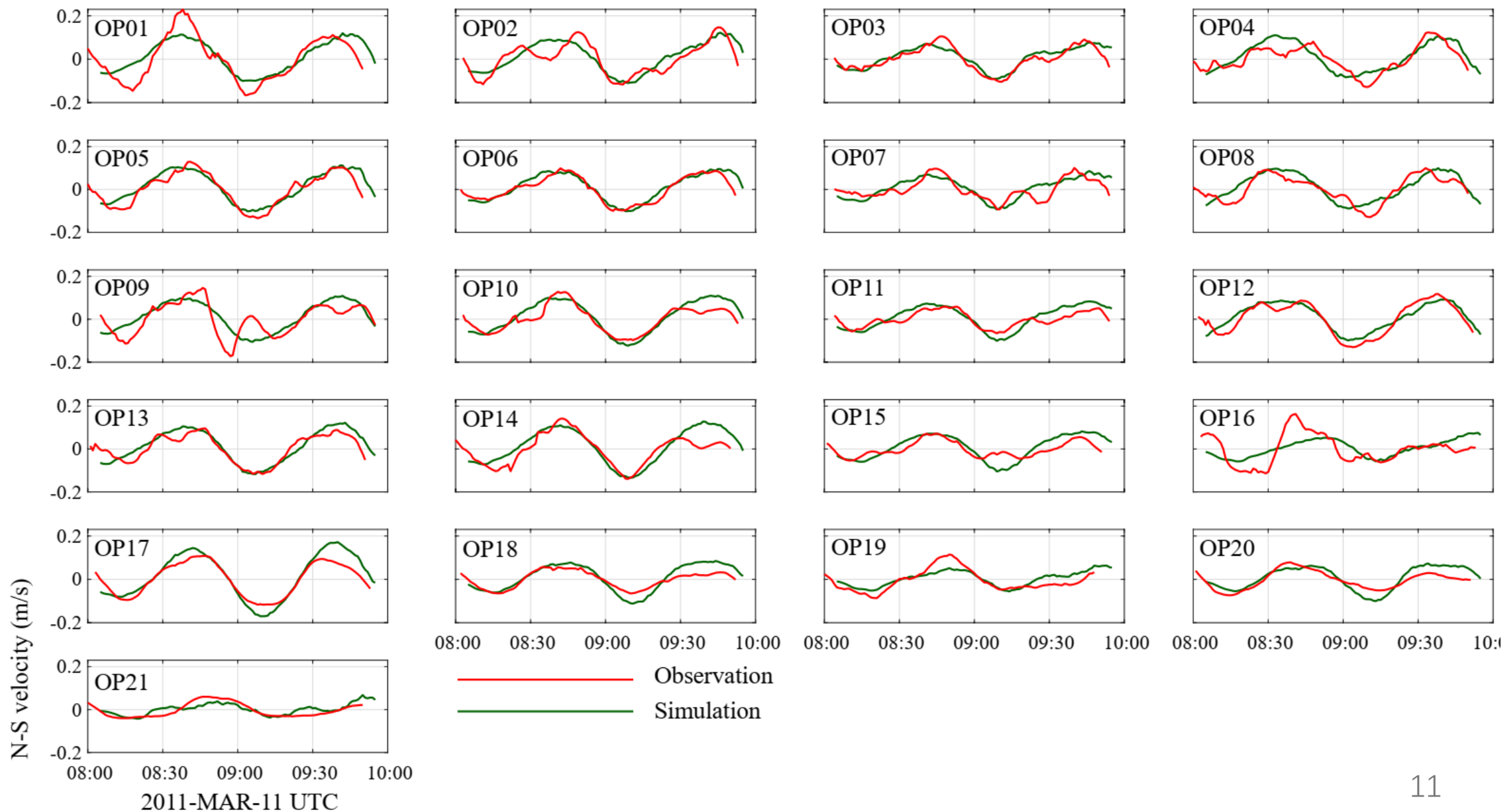
Location of Observational Point (OP)



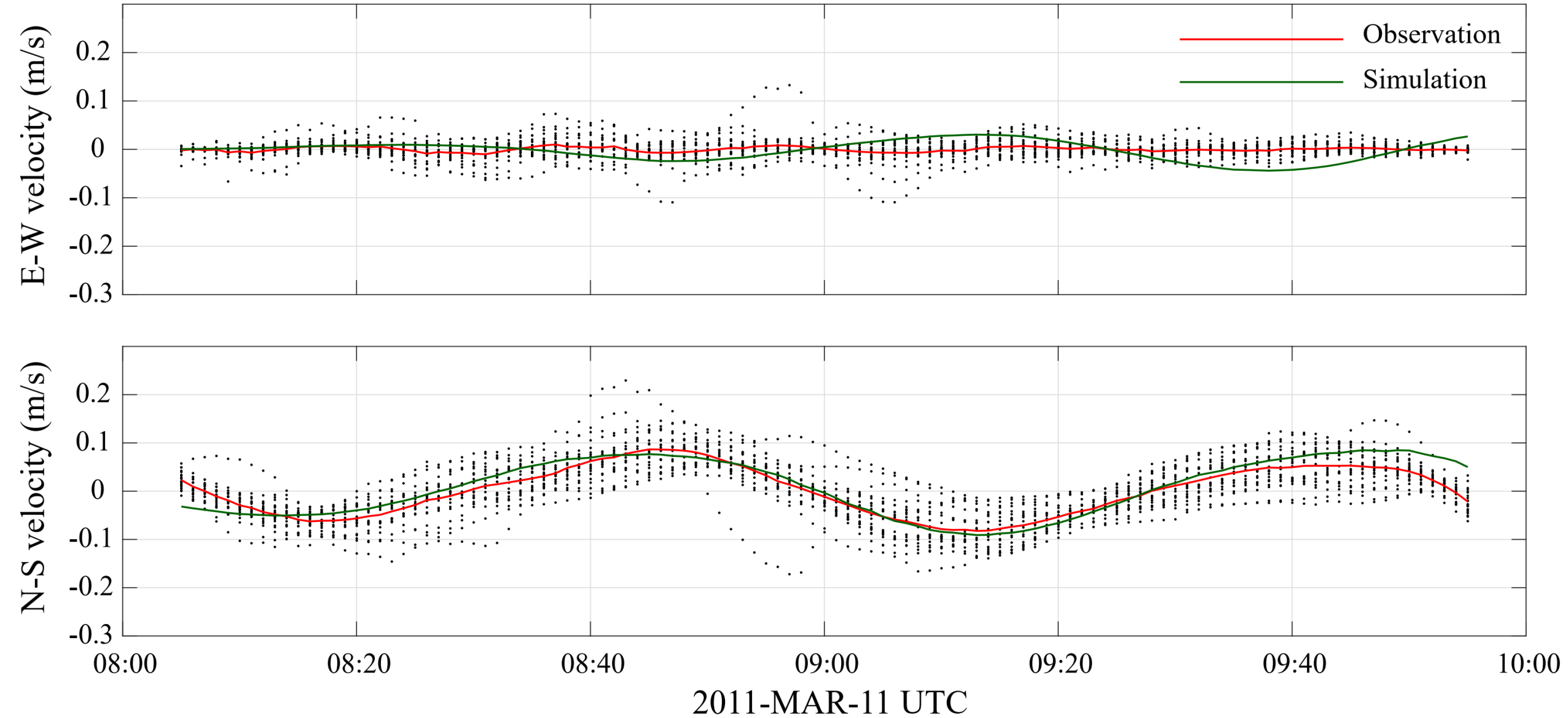
Satake et al. (2013)

- データ解析の正確性を確認するために、**観測された流速**と**シミュレーション値**を比較した。
- シミュレーション値は断層モデルによって計算され、検潮計の波形を正確に再現できることから、流速の再現においても正確であると仮定した。

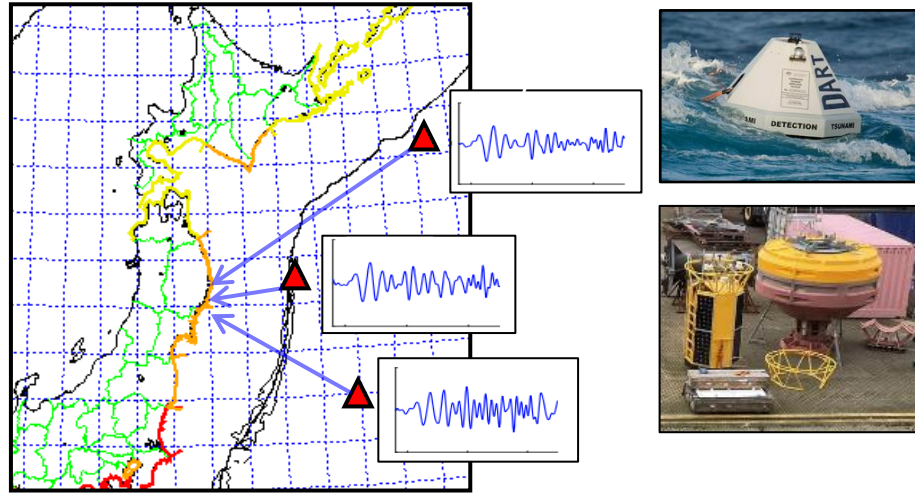
Time Series of Sea Surface Current



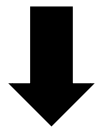
Time Series of Sea Surface Current



Data Assimilation for Early Warning



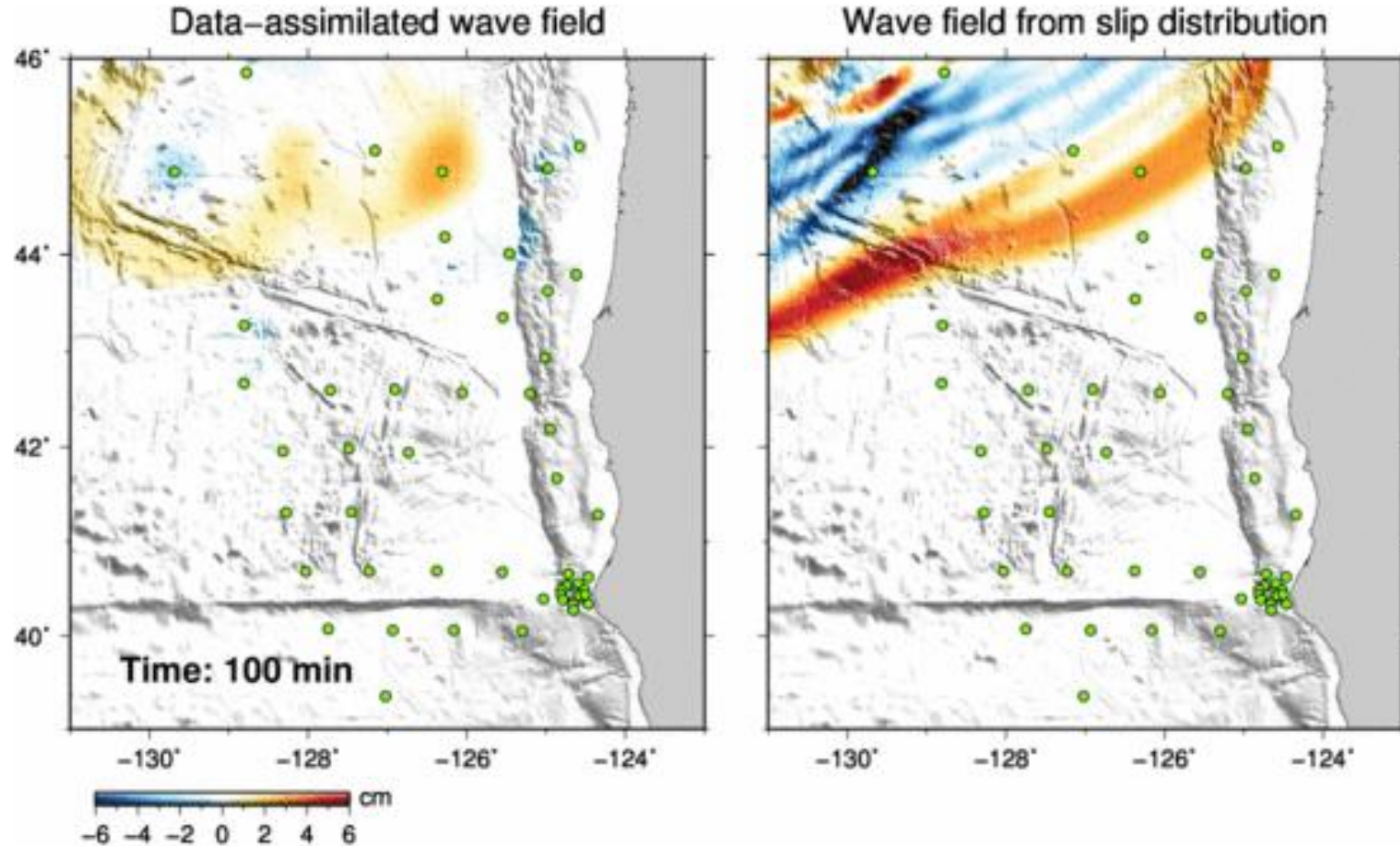
Offshore Tsunami
Observation



Tsunami Forecasting

Do not need source information!

Past study: 2012 Haida Gwaii tsunami (synthetic)



(Gusman et al., [2016](#))

Tsunami wavefield $\mathbf{x}_n = (h(n\Delta t, x, y), P(n\Delta t, x, y), Q(n\Delta t, x, y))^T$

Height X-Velocity Y-Velocity

I. Forecasting Step $\mathbf{x}_n^f(h, P, Q) = \mathbf{F}\mathbf{x}_{n-1}^a(h, P, Q)$

II. Assimilating Step

$$\mathbf{x}_n^a(h, P, Q) = \mathbf{x}_n^f(h, P, Q) + \mathbf{W}(\mathbf{z}_n(\mathbf{P}, \mathbf{Q}) - \mathbf{H}\mathbf{x}_n^f(h, P, Q))$$

$\mathbf{x}_n^f$: Forward tsunami wavefield

$\mathbf{x}_n^a$: Assimilated tsunami wavefield

$\mathbf{F}$: Propagation matrix

$\mathbf{W}$: Weight matrix**

$\mathbf{H}$: Sparse linear observation matrix

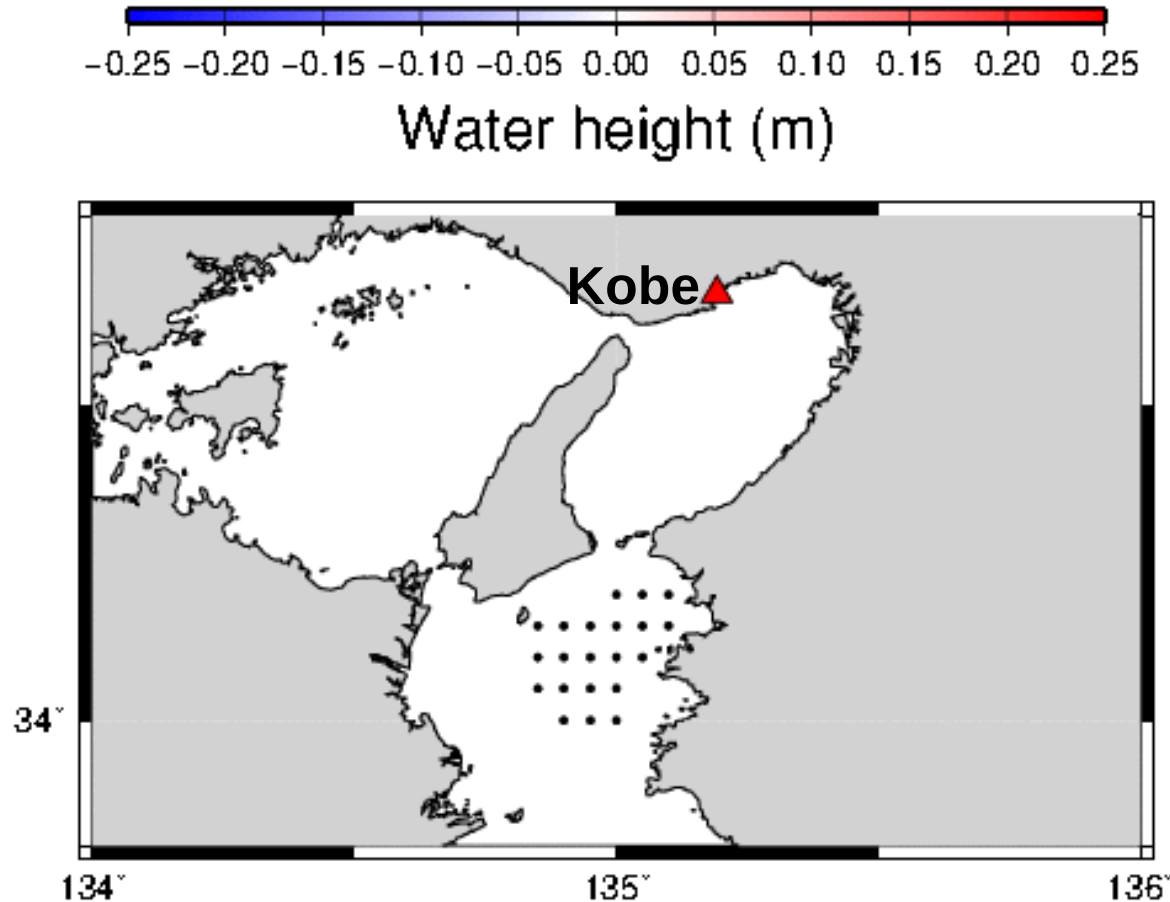
$\mathbf{z}_n$: Observation of HF radar (velocity)

Assimilation starts at 08:05 UTC (JST 17:05)

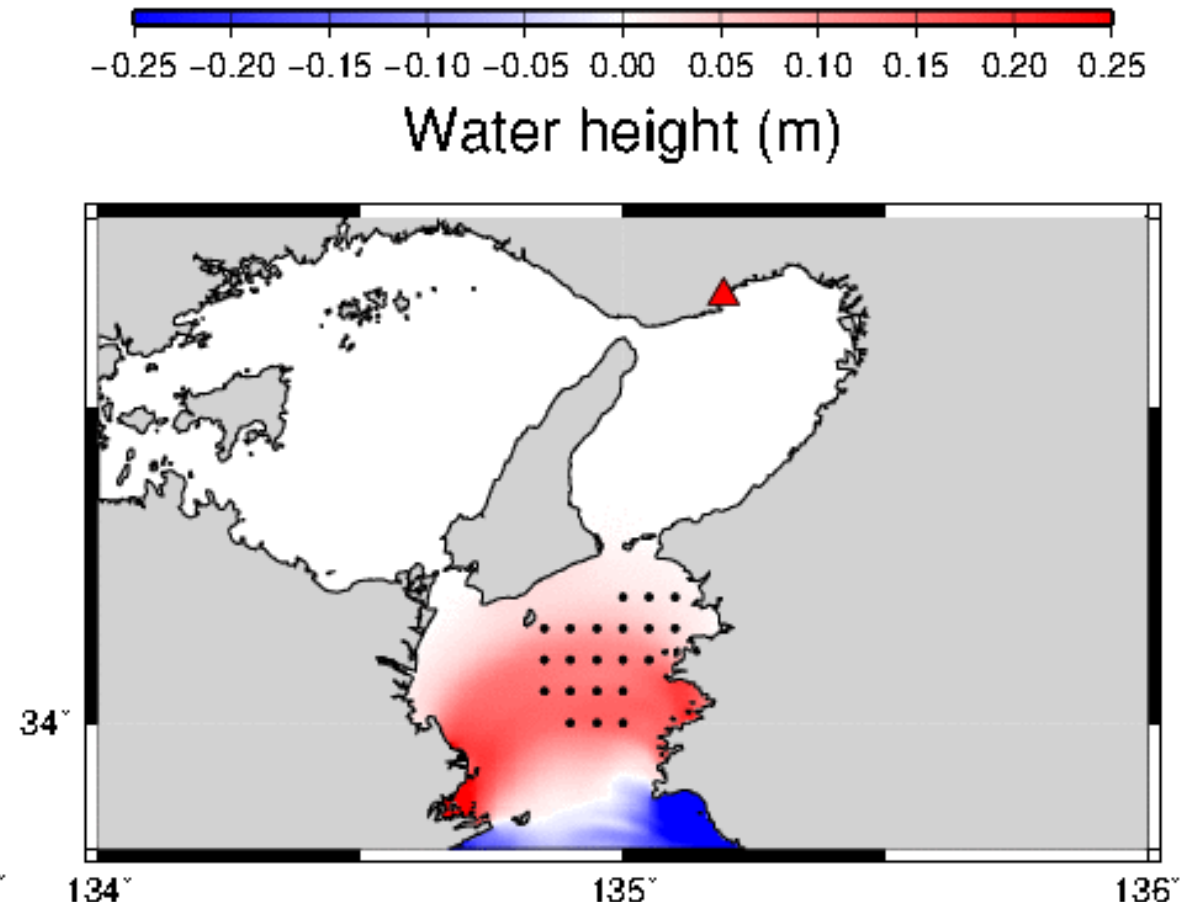
** New algorithm of calculating W was proposed by Dr. Sahana

3. 結果

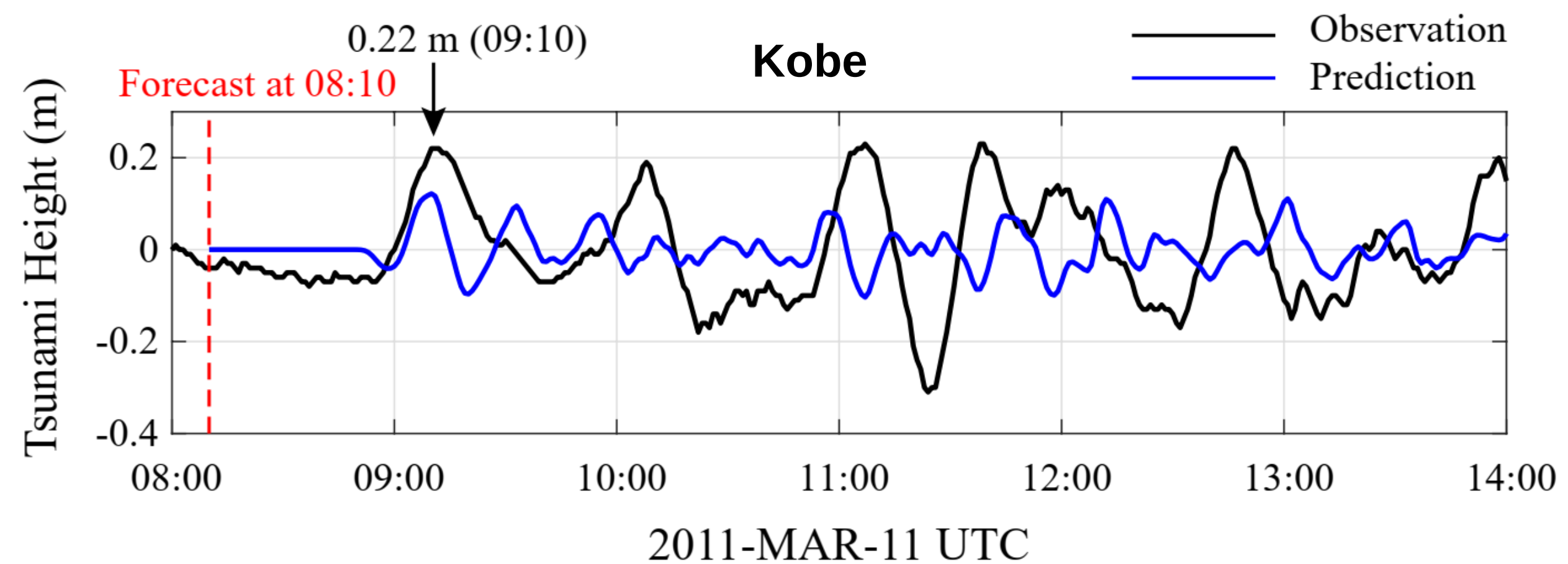
Time: 2011 March 11, 08:05 – 09:00 UTC (17:05 – 18:00 JST)



Forecasted tsunami wavefield using data assimilation of HF radar



Simulated tsunami wavefield computed from source model of Satake et al. (2013)



Observed maximum height: 0.23 m

Predicted maximum height: 0.12 m

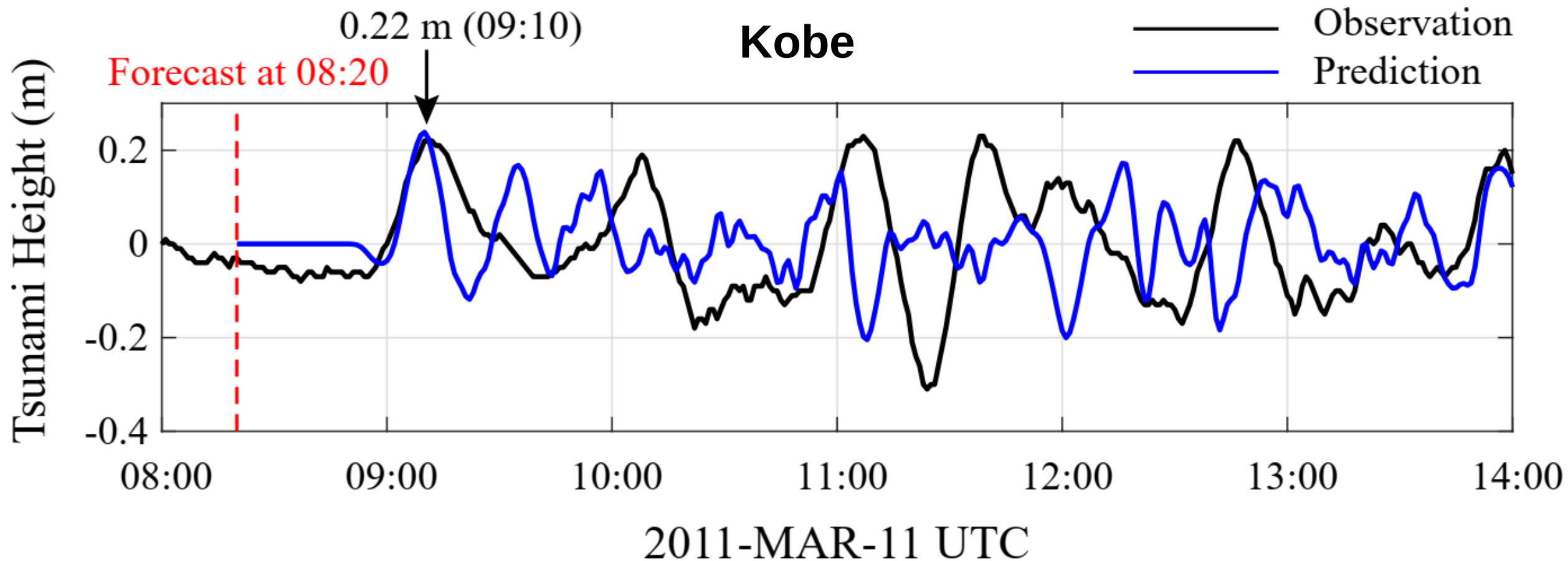
Accuracy (%): 52.2%

Normalized Prediction Error (Navarrete et al., 2019)

$$\min \left\{ \sqrt{\frac{\frac{1}{T} \sum_{t=1}^T (y_{pre}(t+L) - y_{obs}(t))^2}{\frac{1}{T} \sum_{t=1}^T (y_{pre}(t+L))^2 + \frac{1}{T} \sum_{t=1}^T (y_{obs}(t))^2}}; 1 \right\} = 0.991$$

最大振幅：津波早期警報にとって大事

後続波：津波警報解除にとって大事



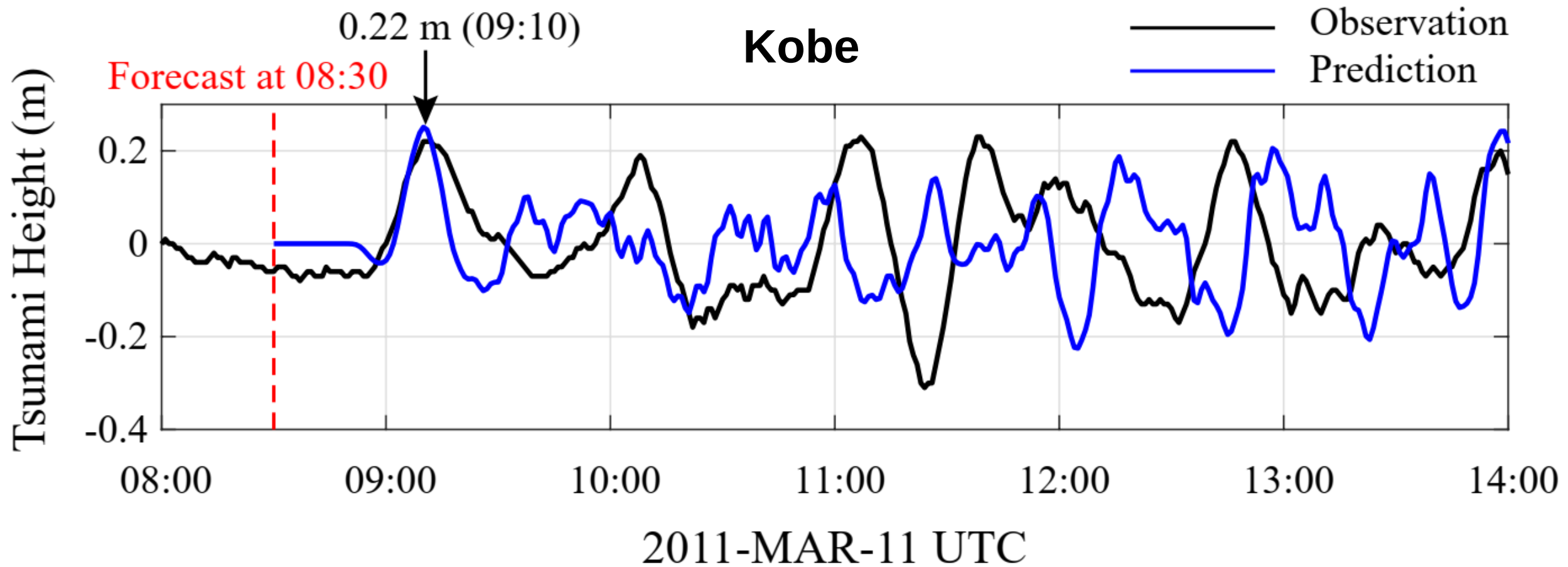
Observed maximum height: 0.23 m

Predicted maximum height: 0.24 m

Accuracy (%): 95.8%

Normalized Prediction Error

$$\min \left\{ \sqrt{\frac{\frac{1}{T} \sum_{t=1}^T (y_{pre}(t+L) - y_{obs}(t))^2}{\frac{1}{T} \sum_{t=1}^T (y_{pre}(t+L))^2 + \frac{1}{T} \sum_{t=1}^T (y_{obs}(t))^2}}; 1 \right\} = 0.976$$



Observed maximum height: 0.23 m

Predicted maximum height: 0.25 m

Accuracy (%): 92.0%

Normalized Prediction Error

$$\min \left\{ \sqrt{\frac{\frac{1}{T} \sum_{t=1}^T (y_{pre}(t+L) - y_{obs}(t))^2}{\frac{1}{T} \sum_{t=1}^T (y_{pre}(t+L))^2 + \frac{1}{T} \sum_{t=1}^T (y_{obs}(t))^2}}; 1 \right\} = 0.929$$

4. まとめ

- HF radar observed the tsunami of the 2011 Tohoku earthquake, with a value of ~ 0.1 m/s at most observational points after 08:00 (UTC).
- Data assimilation of surface current velocity successfully predicted the tsunami waveforms of Kobe tide gauge at 08:20, approximately 50 min prior to tsunami arrival.
- The normalized prediction error decreased over assimilation time.

Thank you for your attention!

Article Navigation

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